



Reassessing the atmospheric oxidation mechanism of toluene

Yuemeng Ji^{a,b,c,1}, Jun Zhao^{c,d,1}, Hajime Terazono^{c,e}, Kentaro Misawa^{c,e,2}, Nicholas P. Levitt^{c,3}, Yixin Li^f, Yun Lin^c, Jianfei Peng^c, Yuan Wang^g, Lian Duan^{c,h}, Bowen Pan^c, Fang Zhang^{c,i}, Xidan Feng^{c,j}, Taicheng An^{a,b,4}, Wilmarie Marrero-Ortiz^f, Jeremiah Secrest^f, Annie L. Zhang^k, Kazuhiko Shibuya^e, Mario J. Molina^{l,4}, and Renyi Zhang^{c,f,4}

^aInstitute of Environmental Health and Pollution Control, Guangdong University of Technology, Guangzhou 510006, China; ^bGuangzhou Key Laboratory of Environmental Catalysis and Pollution Control, School of Environmental Science and Engineering, Guangdong University of Technology, Guangzhou 510006, China; ^cDepartment of Atmospheric Sciences, Texas A&M University, College Station, TX 77843; ^dGuangdong Province Key Laboratory for Climate Change and Natural Disaster Studies, Institute of Earth Climate and Environment System, School of Atmospheric Sciences, Sun Yat-sen University, Guangzhou 510275, China; ^eDepartment of Chemistry, Graduate School of Science and Engineering, Tokyo Institute of Technology, Tokyo 152-8551, Japan; ^fDepartment of Chemistry, Texas A&M University, College Station, TX 77840; ^gDivision of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA 91125; ^hSchool of Resource and Environmental Engineering, East China University of Science and Technology, Shanghai 200237, China; ⁱState Key Laboratory of Earth Surface Processes and Resource Ecology, College of Global Change and Earth System Science, Beijing Normal University, Beijing 100875, China; ^jCollege of Environmental Science and Engineering, Zhongkai University of Agriculture and Engineering, Guangzhou 510225, China; ^kDepartment of Chemistry, College of Natural Sciences, The University of Texas at Austin, Austin, TX 78712; and ^lDepartment of Chemistry and Biochemistry, University of California, San Diego, La Jolla, CA 92093

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Photochemical oxidation of aromatic hydrocarbons leads to tropospheric ozone and secondary organic aerosol (SOA) formation, with profound implications for air quality, human health, and climate. Toluene is the most abundant aromatic compound under urban environments, but its detailed chemical oxidation mechanism remains uncertain. From combined laboratory experiments and quantum chemical calculations, we show a toluene oxidation mechanism that is different from the one adopted in current atmospheric models. Our experimental work indicates a larger-than-expected branching ratio for cresols, but a negligible formation of ring-opening products (e.g., methylglyoxal). Quantum chemical calculations also demonstrate that cresols are much more stable than their corresponding peroxy radicals, and, for the most favorable OH (*ortho*) addition, the pathway of H extraction by O₂ to form the cresol proceeds with a smaller barrier than O₂ addition to form the peroxy radical. Our results reveal that phenolic (rather than peroxy radical) formation represents the dominant pathway for toluene oxidation, highlighting the necessity to reassess its role in ozone and SOA formation in the atmosphere.

aromatics | oxidation | ozone | secondary organic aerosol | air pollution

Toluene is the most abundant aromatic hydrocarbon in the atmosphere and is emitted primarily from anthropogenic sources, i.e., from automobiles and industrial activities. Photochemical oxidation of toluene plays an important role in tropospheric ozone and secondary organic aerosol (SOA) formation, profoundly impacting air quality, human health, and climate (1–8). However, the detailed chemical mechanism of toluene oxidation in the atmosphere remains uncertain (4, 5). Oxidation of toluene is initiated by the hydroxyl radical (OH): the initial OH–toluene reaction results in minor H abstraction (about 10%) and major OH addition (about 90%) (9–20). The H-abstraction pathway leads to the formation of benzaldehyde, whose oxidative pathway is well established (4, 5). The OH addition pathway results in the formation of methylhydroxycyclohexadienyl radicals (the OH–toluene adducts), which subsequently react with O₂ via three plausible pathways (Fig. 1), i.e., H abstraction to yield phenolic compounds and hydroperoxy radicals (HO₂) (pathway I), O₂ addition to form primary peroxy radicals or RO₂ (pathway II), and H abstraction and subsequent O-bridge formation to aromatic oxide/oxepin (pathway III). Previous theoretical studies have suggested that the primary RO₂ cyclizes to form bicyclic radicals, rather than reacting with NO to form alkoxy radicals under atmospheric conditions (16). The bicyclic radicals then undergo unimolecular rearrangement, followed by H abstraction, to form epoxide intermediates or react with O₂ to form secondary RO₂, followed by ring cleavage to produce small α -carbonyl compounds

such as glyoxal and methylglyoxal (16, 21). The toluene–oxide/methylloxepin channel remains speculative, and several previous quantum chemical calculations have shown a high barrier for this pathway (22–26). Numerous experimental studies have been performed to investigate the products from the OH–toluene reactions (2–5, 18, 27). For example, previous studies have reported a highly variable yield of cresols (11, 12, 15, 18, 19, 27), ranging from 9.0 to 52.9%. In addition, the yields of glyoxal and methylglyoxal determined from the previous studies are also conflicting, ranging from <4% to about 39% for glyoxal and <4 to 17% for methylglyoxal (12, 13, 18). From an atmospheric modeling perspective (2–5), ozone

Significance

Aromatic hydrocarbons account for 20 to 30% of volatile organic compounds and contribute importantly to ozone and secondary organic aerosol (SOA) formation in urban environments. The oxidation of toluene, the most abundant aromatic compound, is believed to occur mainly via OH addition, primary organic peroxy radical (RO₂) formation, and ring cleavage, leading to ozone and SOA. From combined experimental and theoretical studies, we show that cresol formation is dominant, while primary RO₂ production is negligible. Our work reveals that the formation and subsequent reactions of cresols regulate the atmospheric impacts of toluene oxidation, suggesting that its representation in current atmospheric models should be reassessed for accurate determination of ozone and SOA formation. The results from our study provide important constraints and guidance for future modeling studies.

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¹Y.J. and J.Z. contributed equally to this work.

²Present address: Department of Chemistry, Graduate School of Science and Engineering, Tokyo Metropolitan University, Tokyo 192-0364, Japan.

³Present address: The Department of Geoscience, University of Wisconsin-Madison, Madison, WI 53706.

⁴To whom correspondence may be addressed. Email: renyi-zhang@tamu.edu, antc99@gdut.edu.cn, or mjmolina@ucsd.edu.

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Table 2. Summary of activation energy (E_a , kcal·mol⁻¹), reaction energy (ΔE_r , kcal·mol⁻¹), rate constant (k , cm³·molecule⁻¹·s⁻¹), and branching ratio (Γ) of the OH–toluene reactions at 298 K

Reactions	Quantity	Ortho	Para	Meta
Toluene + OH → OH adduct	E_a	1.2	2.1	2.7
	ΔE_r	-15.5	-14.6	-14.4
	k	2.7×10^{-12}	7.4×10^{-13}	1.2×10^{-13}
	Γ , %	74*, 80.6 [†] , 52 [‡] , 59	21*, 14.3 [†] , 34 [‡] , 14	3*, 5.1 [†] , 11 [‡] , 5
OH-adduct + O ₂ → cresol	E_a	4.4 [§] , 3.7	4.1	5.0
	ΔE_r	-26.3 [§] , -28.7	-26.8	-27.0
	k	$1.6 \times 10^{-15\ddagger}$, $0.9 \times 10^{-15\ddagger}$, 5.1×10^{-15}	$2.4 \times 10^{-15\ddagger}$, 1.7×10^{-15}	1.2×10^{-16}
OH-adduct + O ₂ → RO ₂	E_a	7.5 [§] , 7.1	0.8	-0.4
	ΔE_r	-8.5 [§] , -2.9	-12.2	-9.5
	k	$1.5 \times 10^{-16\ddagger}$, $2.8 \times 10^{-16\ddagger}$, $1.0 \times 10^{-17\ddagger}$, 3×10^{-15}	$2.5 \times 10^{-14\ddagger}$, 1.5×10^{-14}	4.5×10^{-15}
O-cresol + OH → DHMB	E_a	-0.4		
	ΔE_r	-15.6		
	k	$4.3 \times 10^{-11\#}$, 4.3×10^{-11}		
DHMB + O ₂ → 1,2-dihydroxy-3-ethylbenzene	E_a	2.2		
	ΔE_r	-24.3		
	k	5.4×10^{-12}		

The value is from the present work, except noted otherwise. Γ is calculated by excluding those from the OH *ipso* addition and the H-abstraction pathways.

*From ref. 30.

[†]From ref. 22, including a branching ratio of 3% for OH *ipso* addition.

[‡]From ref. 23, including a branching ratio of 15% for *ipso* position.

[§]From ref. 26.

[¶]From ref. 10.

[#]From ref. 41.

indicating that the secondary products in the OH–toluene reactions undergo additional reactions with OH to form ring-opening products, including small α -carbonyl compounds (glyoxal and methylglyoxal), organic acids, and other low-volatility products (33).

As is evident from Table 2, our calculations of the energetics, rate constants, and branching ratios for the OH–toluene reactions compare favorably with the previously available experimental and theoretical results, considering the respective uncertainties. For example, our calculated rate coefficient of 4.3×10^{-11} cm³·molecule⁻¹·s⁻¹ for OH addition to *o*-cresol to form the DHMB radical is in agreement with the experimental value of $(4.3 \pm 0.5) \times 10^{-11}$ cm³·molecule⁻¹·s⁻¹ by Coeur-Tourneur et al. (41). Additional structural parameters for the reactants, key intermediates, transition states, and products involved in the OH–toluene reaction system are shown in *SI Appendix*, Fig. S5.

Hence, our combined experimental and theoretical methods (42–60) provide kinetic and mechanistic insights into the OH–toluene reactions. In particular, the agreement between our experimental and theoretical results provides compelling evidence for the dominant cresol yield but a negligible formation of methylglyoxal from the initial steps of the OH–toluene reactions. In our work, we quantified only the most abundant peaks detected by our ion drift (ID)-CIMS scheme (using H₃O⁺), accounting for about 64% of the toluene consumed. Another recent experimental study showed that OH addition to the aromatic ring of *o*-cresol leads to hydroxy, dihydroxy, and trihydroxy methyl benzoquinones and dihydroxy, trihydroxy, tetrahydroxy, and pentahydroxy toluenes, detected in the gas phase by CIMS (using CF₃O⁻ and H⁺·*n*H₂O where *n* = 1, 2, ...) and in the particle phase using offline direct analysis with real-time mass spectrometry (33).

The differences in the measured product yields between our present experimental work and the previous laboratory experiments are explainable by the distinct conditions among the various experimental studies. Although the environmental chamber method has been applied extensively in development of parameterizations of formation of ozone and SOA for atmospheric modeling purposes (5), there were several intricate difficulties for the chamber approach, which made it rather unsuitable for detailed kinetic and mechanistic investigations of atmospheric

hydrocarbon chemistry. Specifically, the limitations in the earlier chamber studies included longer reaction times (minutes to hours), higher reactant concentrations, wall loss, and the lack of online detection and quantification of reactive reactants and products by advanced analytical instruments. For example, an earlier experimental study reported a yield of 25.2% for the cresols from the OH–toluene reactions (27). The concentrations of the reactants used in that study were about 4×10^{13} and 1×10^{13} molecules·cm⁻³ for toluene and NO_x, respectively, and the OH concentration was not measured but was estimated from the decay in the toluene concentration. The concentrations of the reactants and products were measured by using offline gas chromatography with flame ionization detection. For the irradiation time of about 10 min during their chamber experiments, a lifetime of cresols was estimated to be about 230 s, using the rate coefficient of 4.3×10^{-11} cm³·molecule⁻¹·s⁻¹ for the reaction between cresols and OH from this work and another experimental study (40). Hence, there likely existed significant secondary reactions of cresols with OH in the earlier chamber investigations, responsible for the measured lower cresol yields in the majority of those studies (12, 19, 27).

On the other hand, the technique of high-pressure turbulent flow reactors in conjunction with CIMS detection has been developed for accurate kinetic measurements of atmospheric gas-phase reactions (35, 36). The main advantage of the fast-flow reactor system lies in the ability to isolate the individual reaction steps and intermediates (35, 36, 42–44, 48, 61). Specifically, the CIMS technique allows for online detection of many reactants, intermediates, and products with high sensitivity and selectivity (35, 36). In addition, the ID-CIMS method provides quantification of the gas-phase concentrations of the intermediates and products without the necessity of calibration, which is advantageous because of the general difficulty in obtaining the authentic standards for products of hydrocarbon reactions (34–36). Furthermore, a turbulent flow condition effectively minimizes the wall loss (35, 36). In our experiments, the reactant concentrations were $(0.4$ to $1.6) \times 10^{12}$ and $(5.4$ to $19) \times 10^{12}$ molecules·cm⁻³ for toluene and NO₂, respectively, about one to two orders of magnitude lower than those of the environmental chamber studies (12, 19, 27). In addition, the OH concentration was directly quantified in our experiments, in a range of $(0.8$ to $8.0) \times 10^{10}$ molecules·cm⁻³. Secondary reactions were

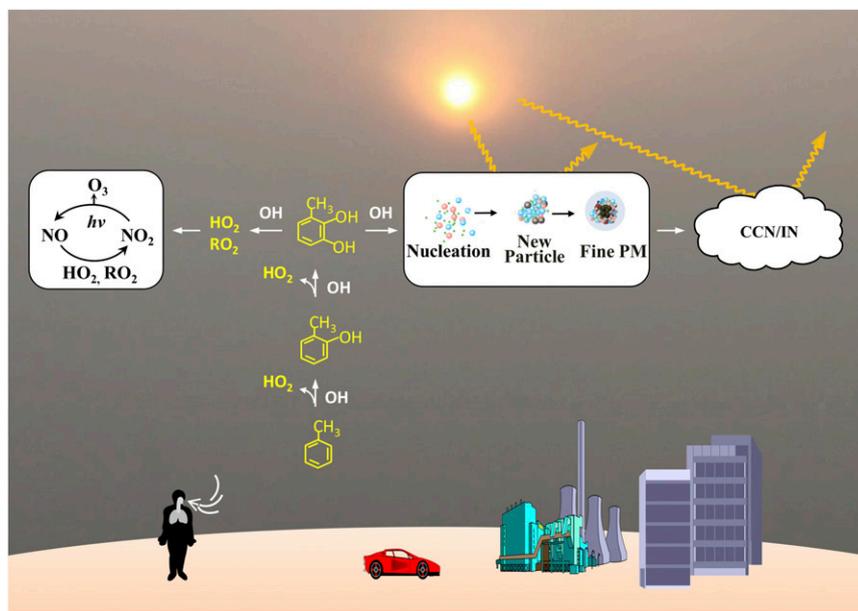


Fig. 4. Toluene oxidation and urban photochemical smog. The OH–toluene reactions lead promptly and dominantly to cresol formation, which regulates the impacts on air quality, human health, and climate. The phenolic pathway yields HO₂ and RO₂ that facilitate ozone production by converting NO to NO₂, as photodissociation of NO₂ occurs in the wavelength range of 400 nm to 650 nm (the symbol $h\nu$ denotes the energy of a photon, where h is the Planck number and ν is the frequency of light). Also, dihydroxymethylbenzenes generated from this pathway undergo subsequent reactions with OH to form the precursors for aerosol nucleation and growth, including small α -carbonyl compounds, organic acids, and other highly oxygenated low-volatility products.

effectively suppressed in our experiments because of the short reaction time in the flow reactor (on the order of about 50 ms). The lifetime of cresols was estimated to be about 2 s in our experiments, responsible for our measured high yield of cresols. Under our experimental conditions, the self-reaction of the OH–toluene adducts was unimportant on the basis of this rate constant reported previously (10); this reaction was estimated to be about one to two orders of magnitude lower than that of the OH–toluene adduct with O₂. Furthermore, the absolute O₂ concentration had no effect for the competing reactions between the cresol and primary RO₂ pathways (i.e., pathway I vs. pathway II in Fig. 1), and the previous experimental results did not show a pressure dependence of the OH–toluene reaction system (61).

Also, our theoretical calculations indicate that formation of glyoxal and methylglyoxal via primary RO₂ is minimal, consistent with the absence in the ring-opening α -carbonyl products detected in our experiments. It is anticipated that those species are formed as the multigeneration products from the cresol pathway (33), likely similar to the cases with their high yields in the previous chamber experiments (12, 13).

Conclusion

Photochemical oxidation of toluene contributes importantly to ozone and SOA formation in urban environments, with key implications for air quality and human health (1–3, 62). In addition, by directly scattering and absorbing solar radiation and indirectly serving as cloud condensation nuclei or ice nuclei, SOA represents the major component in global radiative forcing on climate (1–3). For toluene oxidation, ozone formation in current atmospheric models is represented mainly via primary RO₂, with a minor contribution from cresols (5). In contrast, our experimental and theoretical results reveal the exclusive prompt productions of cresols and HO₂, but insignificant formation of primary RO₂ from the reaction of OH–toluene adducts with O₂. HO₂ is also produced from the reaction of cresols with OH, and further reactions of dihydroxymethylbenzenes with OH yield additional HO₂ and RO₂. As a result, the production of HO₂ and RO₂ along the phenolic pathway regulates the abundance of the total peroxy radicals, which facilitates the conversion between NO and NO₂ (Fig. 4). In addition, the relative branching between the RO₂ and cresol pathways likely affects the OH radical propagation and termination, which are also important for ozone formation.

The further oxidation of dihydroxymethylbenzenes by OH leads to formation of small α -dicarbonyl compounds (glyoxal and methylglyoxal), organic acids, and other low-volatility products

(33, 40), which likely contribute to aerosol nucleation and growth by partitioning and particle-phase reactions (1, 63–65). In particular, the recent experimental study by Schwantes et al. (33) has identified a substantial fraction of highly oxygenated low-volatility products, contributing to 20 to 40% of SOA formation from the cresol oxidation initiated by OH radicals.

Our combined experimental and theoretical results show that the prompt formation of cresols and their subsequent oxidation largely regulate the atmospheric impacts of toluene oxidation (Fig. 4), indicating that the representation of the toluene oxidation mechanism should be reassessed in current atmospheric models. Future studies are necessary to incorporate the kinetic and mechanistic results from our present work into atmospheric models, to accurately assess ozone and SOA formation under polluted environments.

Materials and Methods

The experiments were performed using a fast-flow reactor in conjunction with ID-CIMS. Detailed description of the ID-CIMS technique can be found in our previous publications (42–44).

Quantum chemical calculations were performed with the Gaussian 09 suite of programs (45). Geometry optimization of the relevant species were executed at the M06-2X level with the standard 6-311G(d,p) basis set [M06-2X/6-311G(d,p)]. This level of theory has been successfully applied to atmospheric reactions (46, 47). The vibrational analysis was made at the same level of theory to characterize the nature of each critical point along the potential energy surface (PES) with a local minimum or a transition state (exactly one imaginary frequency) and to make zero-point-energy corrections. The minimum-energy path was constructed with the intrinsic reaction coordinate theory to confirm that the transition state connected with the minima along the reaction path. Because kinetic calculations of the organic reaction systems were highly sensitive to the predicted energetics (48–51), single-point energy calculations were performed to refine the PES using the QCISD(T)/6-311+G(2d,p) level. The dual-level potential profile along the reaction path was further refined with the interpolated single-point energy method (52), in which extra single-point calculations were performed to correct the lower-level reaction path. The dual-level dynamics approach was denoted as X/Y, where a single-point energy calculation at level X was carried out for the geometry optimized at a lower level Y. Using the Polyrate 2010A (53) and KiThelp programs (54), the rate constants were calculated according to the canonical variational transition state theory (55–59), along with the small curvature tunneling correction (60).

Additional descriptions of the experimental conditions and the structure of the reactants, products, intermediates, and transition states are provided in *SI Appendix*.

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